

Challenges in Inverse Modeling and Data Assimilation of Atmospheric Constituents

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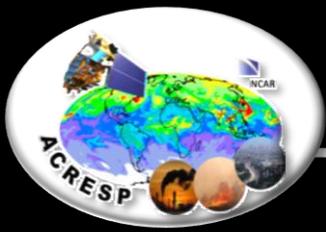
National Center for Atmospheric Research

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with input from **Monika Kopacz**

Woodrow Wilson School of International and Public Affairs

Princeton University

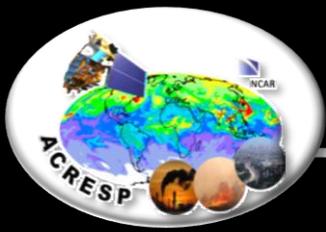


the chemical data assimilation problem

*Our goal is to combine **observations** and **models** to improve our estimate of the **state** (**constituents**) of a system (atmosphere).*

from a Bayesian framework:

$$\begin{array}{l} \text{(state given observations)} \propto \text{(observations given state)} \text{(state)} \\ [\mathbf{x} | \mathbf{Y}] \qquad \alpha \quad [\mathbf{Y} | \mathbf{x}] \qquad [\mathbf{x}] \\ \text{(posterior)} \qquad \alpha \quad \text{(likelihood)} \qquad \text{(prior)} \end{array}$$



the inverse problem

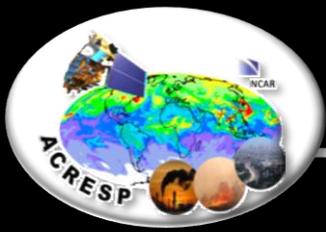
*Our goal is to combine **observations** and **models** to improve our estimate of the **sources** (of constituents) in the system (atmosphere).*

from a Bayesian framework:

(source given observations of the state)

\propto (observations given the source) (source)

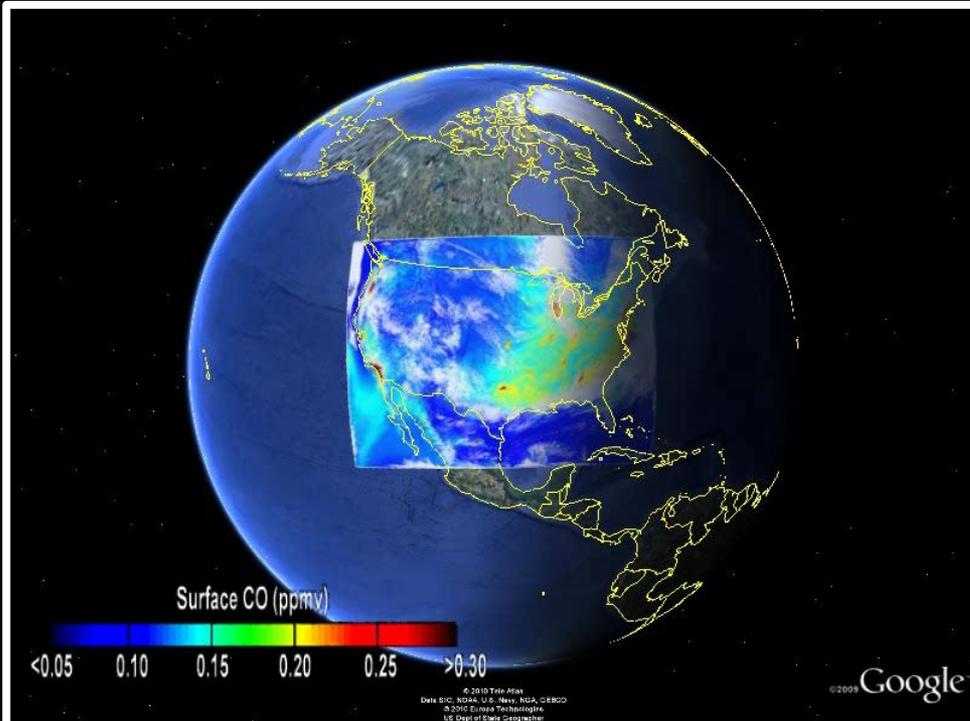
$$\begin{array}{l} [\mathbf{x} | \mathbf{Y}] \\ \text{(posterior)} \end{array} \propto \begin{array}{l} [\mathbf{Y} | \mathbf{x}] \\ \text{(likelihood)} \end{array} \quad \begin{array}{l} [\mathbf{x}] \\ \text{(prior)} \end{array}$$



our model of the system

estimate = $M(\text{state}) + \text{error}$

$$\mathbf{x}_{t+1} = \mathbf{M}(\mathbf{x}_t) + \eta_t \text{ with } \eta_t \sim N(0, Q)$$



from a 4km WRF-Chem
North American Monsoon (NAM) Simulation
(~July/Aug 2006)

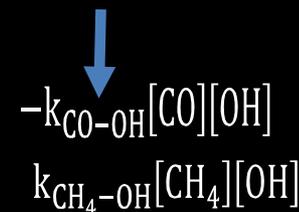
P.I.s Mary Barth/Alma Hodzic

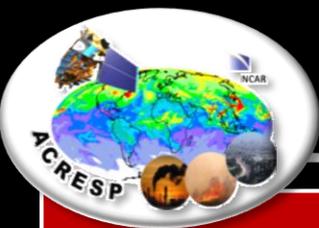
<http://acd.ucar.edu/~barthm/namcase.html>

In the case of [CO]:

$$\frac{d[\text{CO}]}{dt} = \left(\frac{\partial[\text{CO}]}{\partial t} \right)_{\text{transport}} + \left(\frac{\partial[\text{CO}]}{\partial t} \right)_{\text{emissions}} + \left(\frac{\partial[\text{CO}]}{\partial t} \right)_{\text{chemistry}} + \left(\frac{\partial[\text{CO}]}{\partial t} \right)_{\text{deposition}}$$

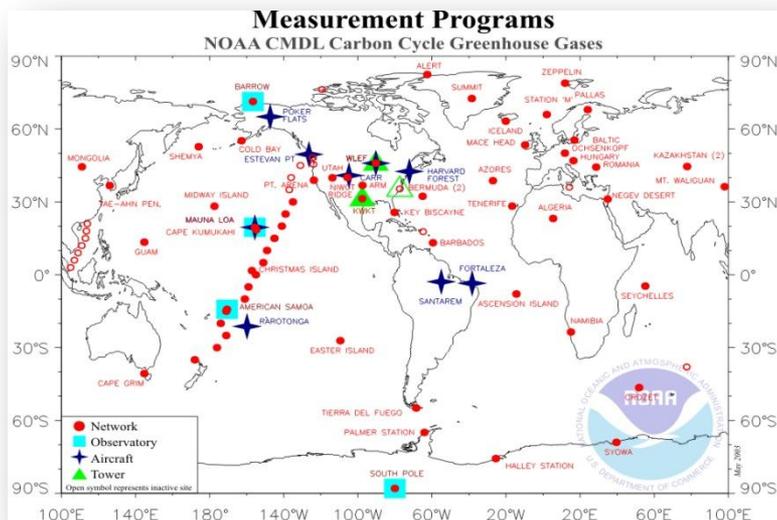
$-\vec{v} \cdot \nabla[\text{CO}]$ (transport) natural / anthropogenic combustion-related processes (emissions)



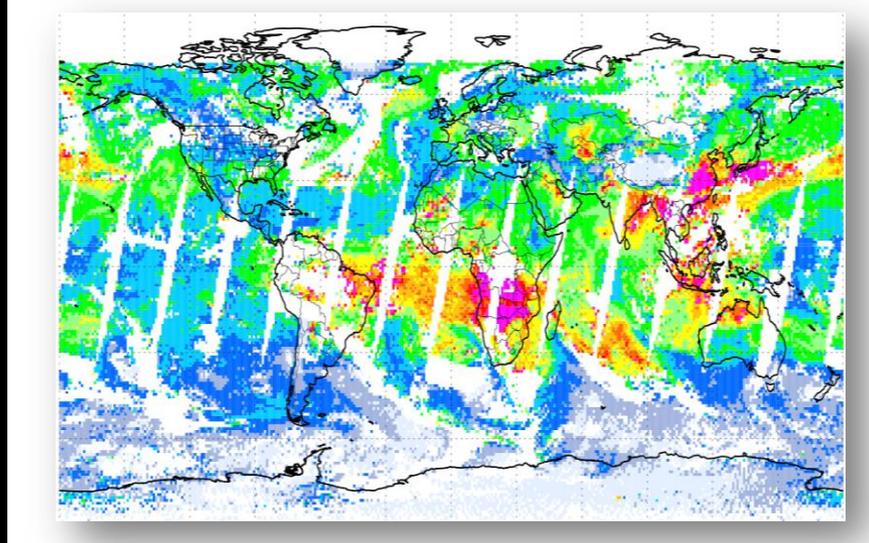


observations of the system

In-Situ Measurements

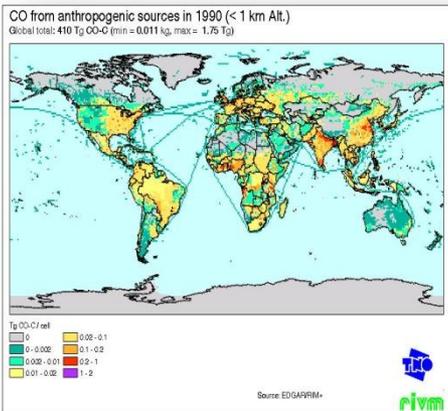


Remote-sensed Measurements



Ancillary Data

e.g. Terra MOPITT, Aura TES, Envisat SCIAMACHY, Aqua AIRS, Aura MLS, ACE FTS, Metop IASI

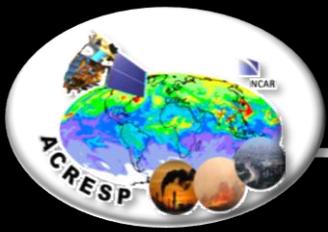


Aircraft Measurements



Source Inventories

Fire Data



observation and model

DA problem

observation = $h(\text{state}) + \text{error}$

$$\mathbf{Y} = \mathbf{h}(\mathbf{x}) + \boldsymbol{\varepsilon} \quad \text{with } \boldsymbol{\varepsilon} \sim N(\mathbf{0}, \mathbf{R})$$

$$\mathbf{Y} = \mathbf{H}\mathbf{x} + \boldsymbol{\varepsilon} \quad \text{linear case}$$

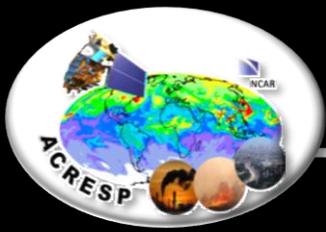
In the case of direct observations of CO, \mathbf{H} is simply a linear interpolation of the model CO state to the observation

However, in the case of remote-sensed measurements of CO, \mathbf{H} can be more complicated (may not be non-linear too!). For example,

$\mathbf{y}_{\text{radiance}} = \mathbf{h}(\mathbf{x}) + \mathbf{e}_y$ where $\mathbf{h}(\cdot)$ can be a radiative transfer model (RTM)

$\mathbf{y}_{\text{radiance}} = \frac{\delta \mathbf{y}}{\delta \mathbf{x}} \mathbf{x} + \mathbf{e}_y$ linear case where $\frac{\delta \mathbf{y}}{\delta \mathbf{x}}$ is the measurement sensitivity

$\mathbf{y}_{\text{retrieval}} = \hat{\mathbf{x}} = \mathbf{x}_a + \frac{\delta \hat{\mathbf{x}}}{\delta \mathbf{x}} (\mathbf{x} - \mathbf{x}_a) + \mathbf{e}_x$ where $\mathbf{A} = \frac{\delta \hat{\mathbf{x}}}{\delta \mathbf{x}}$ is the averaging kernel



observation and model

Inverse problem

observation = $h(\text{source}) + \text{error}$

$$\mathbf{Y} = \mathbf{h}(\mathbf{x}) + \boldsymbol{\varepsilon} \quad \text{with } \boldsymbol{\varepsilon} \sim N(\mathbf{0}, \mathbf{R})$$

$$\mathbf{Y} = \mathbf{H}\mathbf{x} + \boldsymbol{\varepsilon} \quad \text{linear case}$$

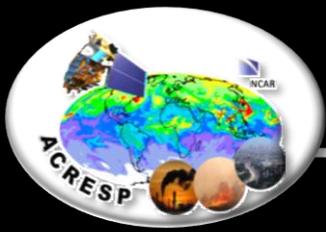
In the case of direct observations of CO, \mathbf{H} is matrix of response functions (mapping the source to the state) + a linear interpolation of the model CO state to the observation

However, in the case of remote-sensed measurements of CO, \mathbf{H} can be more complicated (may not be non-linear too!). For example,

$$\mathbf{y}_{\text{radiance}} = \mathbf{h}(\mathbf{x}) + \mathbf{e}_y \quad \text{where } \mathbf{h}(\cdot) \text{ can be a radiative transfer model (RTM)}$$

$$\mathbf{y}_{\text{radiance}} = \frac{\delta \mathbf{y}}{\delta \mathbf{x}} \mathbf{x} + \mathbf{e}_y \quad \text{linear case where } \frac{\delta \mathbf{y}}{\delta \mathbf{x}} \text{ is the measurement sensitivity}$$

$$\mathbf{y}_{\text{retrieval}} = \hat{\mathbf{x}} = \mathbf{x}_a + \frac{\delta \hat{\mathbf{x}}}{\delta \mathbf{x}} (\mathbf{x} - \mathbf{x}_a) + \mathbf{e}_x \quad \text{where } \mathbf{A} = \frac{\delta \hat{\mathbf{x}}}{\delta \mathbf{x}} \text{ is the averaging kernel}$$



the inverse problem

$[\mathbf{x} | \mathbf{Y}]$
(posterior)

$[\mathbf{Y} | \mathbf{x}]$
(likelihood)

$[\mathbf{x}]$
(prior)

is $N(\mathbf{x}^f, \mathbf{P}^f)$ and $\mathbf{Y} = \mathbf{H}\mathbf{x} + \varepsilon$ with $\varepsilon \sim N(0, \mathbf{R})$, the posterior is $N(\mathbf{x}^a, \mathbf{P}^a)$

$$\mathbf{J}_{\mathbf{x}} = (\mathbf{y} - \mathbf{H}\mathbf{x}^f)^t \mathbf{R}^{-1} (\mathbf{y} - \mathbf{H}\mathbf{x}^f) + (\mathbf{x} - \mathbf{x}^f)^t \mathbf{P}^{f-1} (\mathbf{x} - \mathbf{x}^f)$$

An estimate \mathbf{x}^a of the state can be expressed as:

$$\mathbf{x}^a = \left(\mathbf{H}^t \mathbf{R}^{-1} \mathbf{H} + \mathbf{P}^{f-1} \right)^{-1} \left(\mathbf{H}^t \mathbf{R}^{-1} \mathbf{y} + \mathbf{P}^{f-1} \mathbf{x}^f \right), \quad \mathbf{P}^a = \left(\mathbf{H}^t \mathbf{R}^{-1} \mathbf{H} + \mathbf{P}^{f-1} \right)^{-1}$$

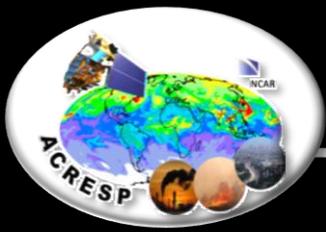
Kalman Filter

$$\mathbf{x}^a = \mathbf{x}^f + \mathbf{K}(\mathbf{y} - \mathbf{H}\mathbf{x}^f), \quad \mathbf{P}^a = (\mathbf{I} - \mathbf{K}\mathbf{H})\mathbf{P}^f$$

$$\mathbf{x}^a = \mathbf{x}^f + \mathbf{P}^f \mathbf{H}^t (\mathbf{H}\mathbf{P}^f \mathbf{H}^t + \mathbf{R})^{-1} (\mathbf{y} - \mathbf{H}\mathbf{x}^f)$$

Can be recast as:

$$\mathbf{x}^a = \underbrace{\mathbf{x}^f + \mathbf{P}^f \mathbf{H}^t (\mathbf{H}\mathbf{P}^f \mathbf{H}^t)^{-1}}_{\text{least-squares}} \underbrace{(\mathbf{H}\mathbf{P}^f \mathbf{H}^t) (\mathbf{H}\mathbf{P}^f \mathbf{H}^t + \mathbf{R})^{-1}}_{\text{shrink to } \mathbf{H}\mathbf{x}^f} (\mathbf{y} - \mathbf{H}\mathbf{x}^f)$$



inverse modeling of CO sources

e.g. $[\text{emission} \mid \text{MOPITT}] \sim [\text{MOPITT} \mid \text{emission}] [\text{emission}]$

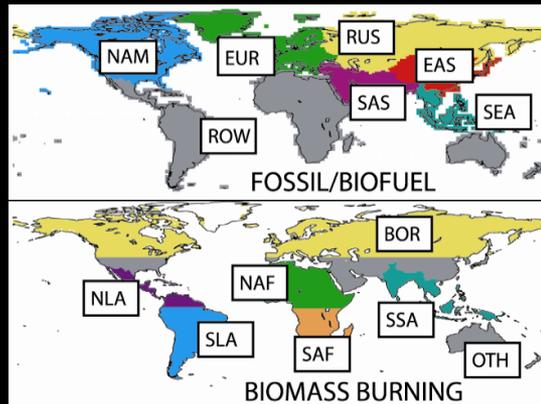
$N(\mathbf{x}^a, \mathbf{P}^a)$

 $N(\mathbf{H}\mathbf{x}, \mathbf{R})$

 $N(\mathbf{x}^f, \mathbf{P}^f)$

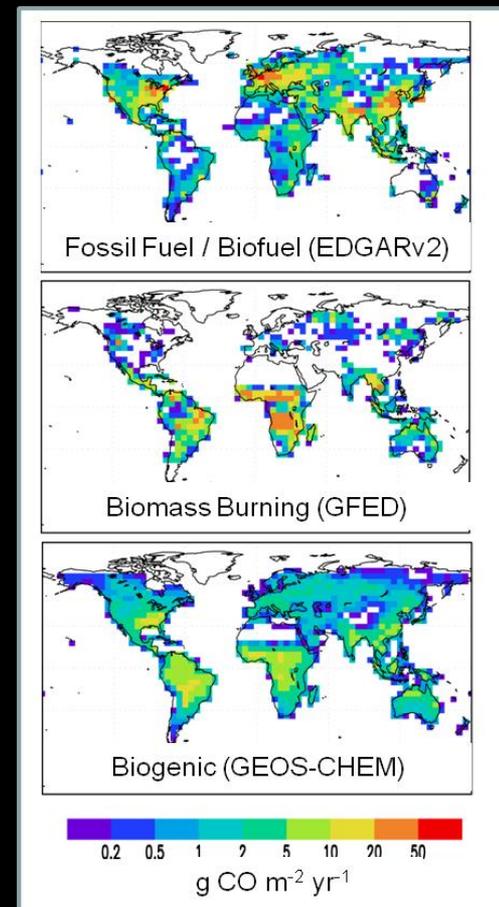
We solve for regional source scaling factors

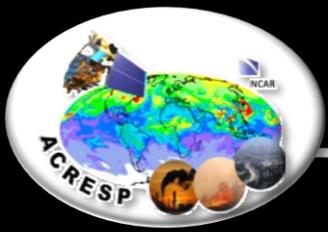
e.g.



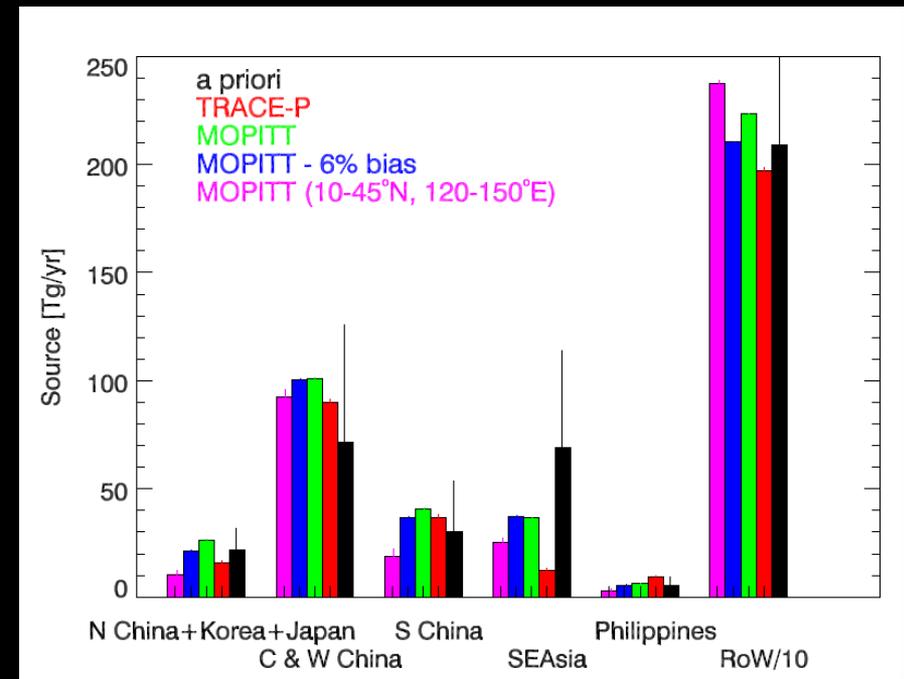
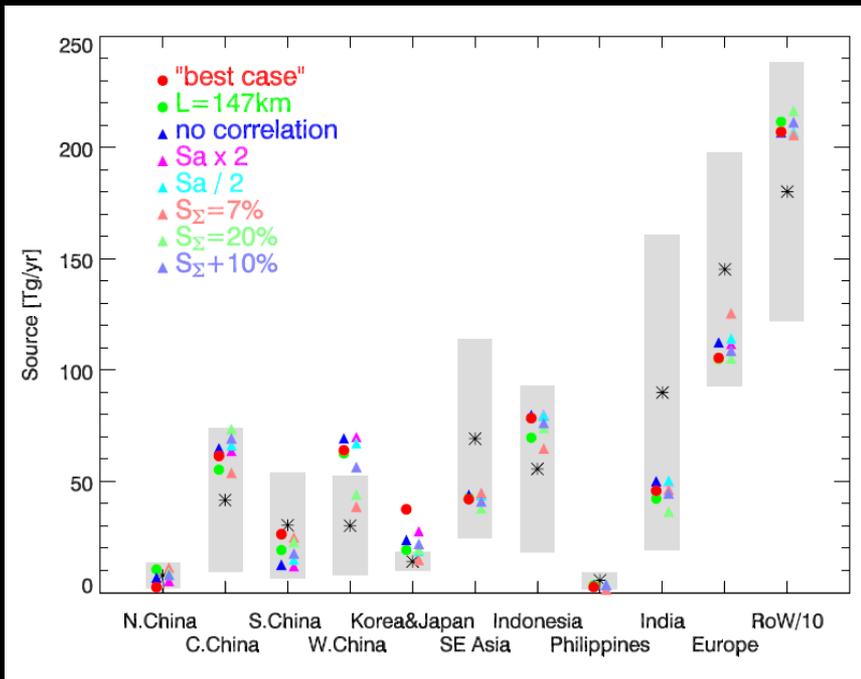
Note that the posterior mean is sensitive to data, prior estimates, error specification and model leading to persistent discrepancies in source estimates

(No TransCom framework for CO)

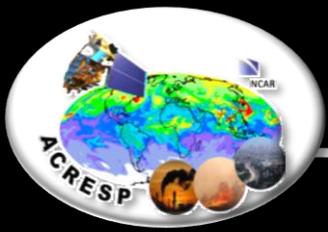




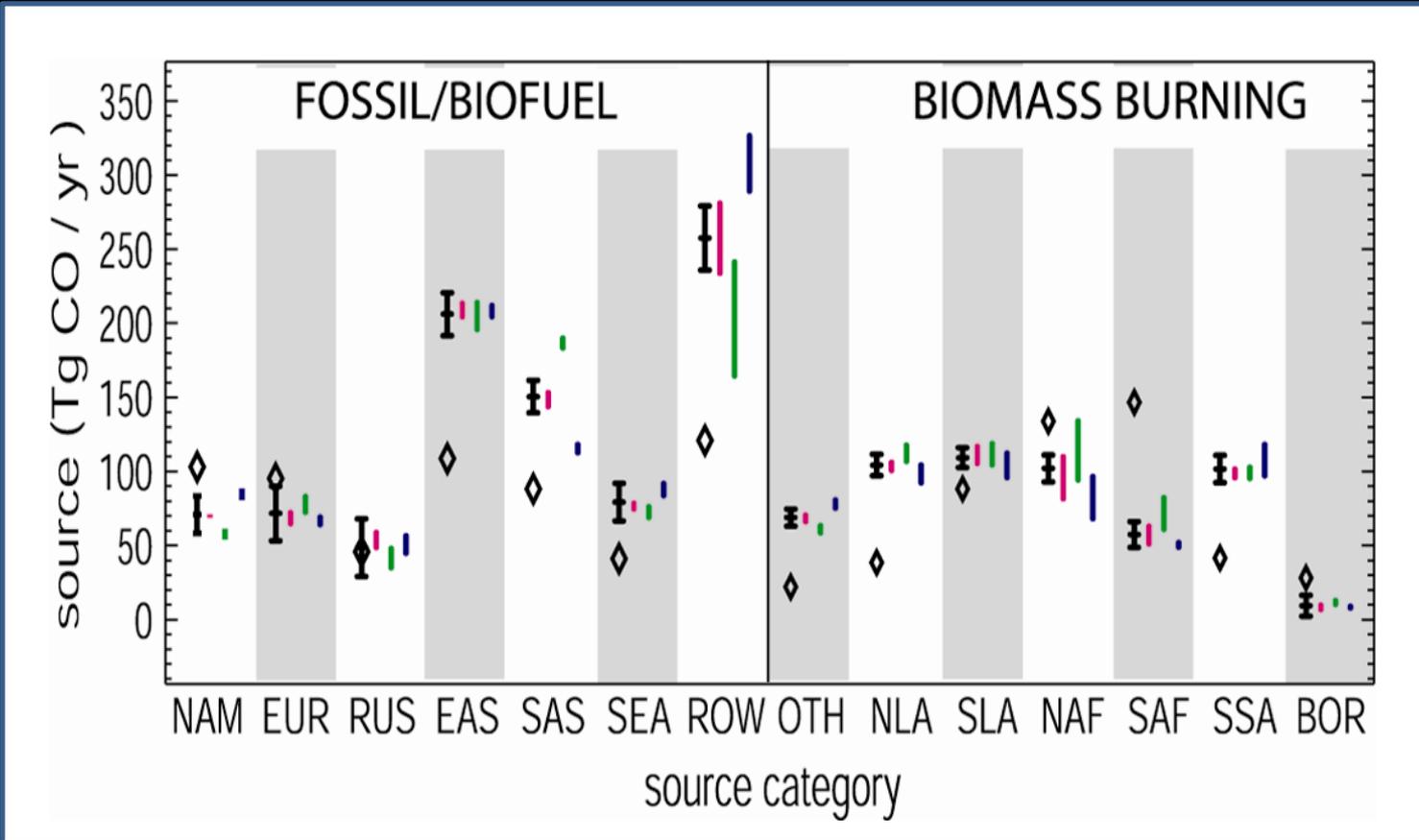
Sensitivity Tests of Error Assumptions and Obs Data Choice



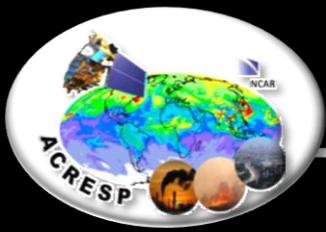
Heald et al. (2004)



Sensitivity Tests of Error Assumptions and Obs Data Choice

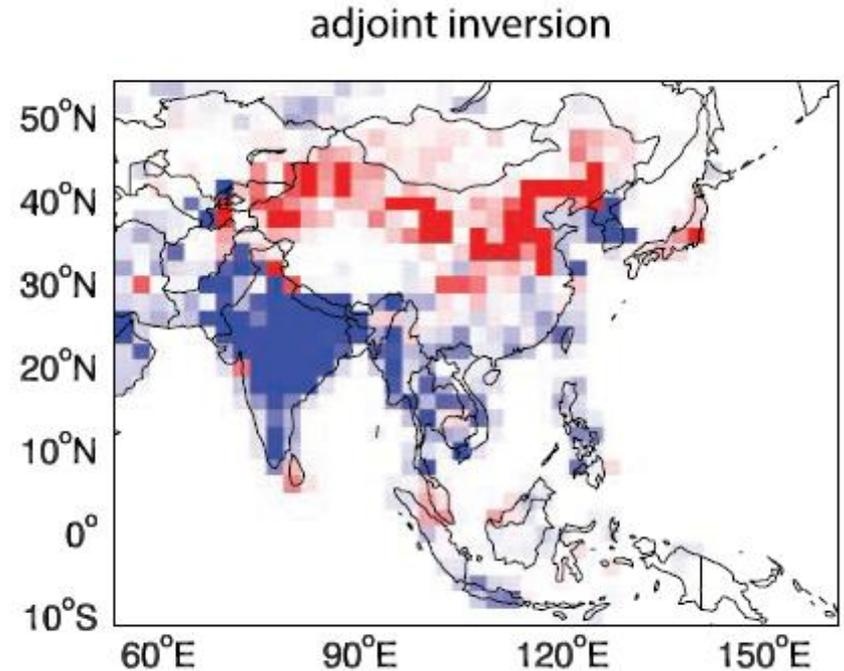
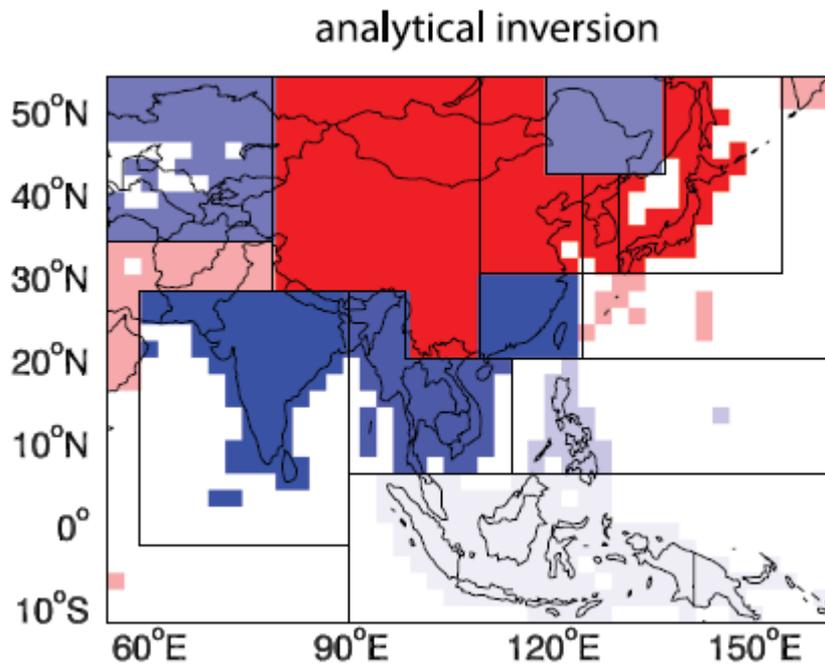


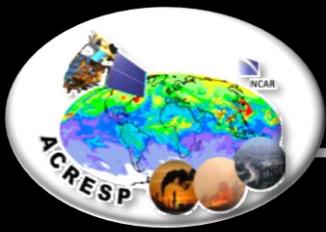
Arellano et al. (2004)



sensitivity to methodology

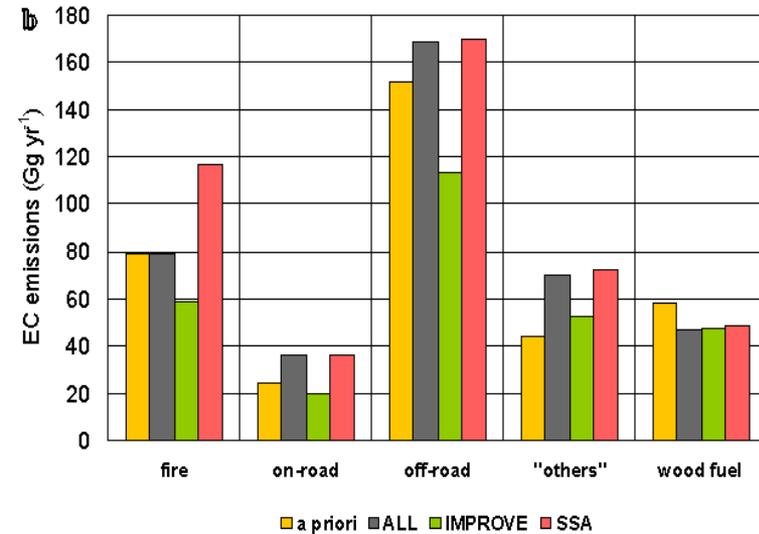
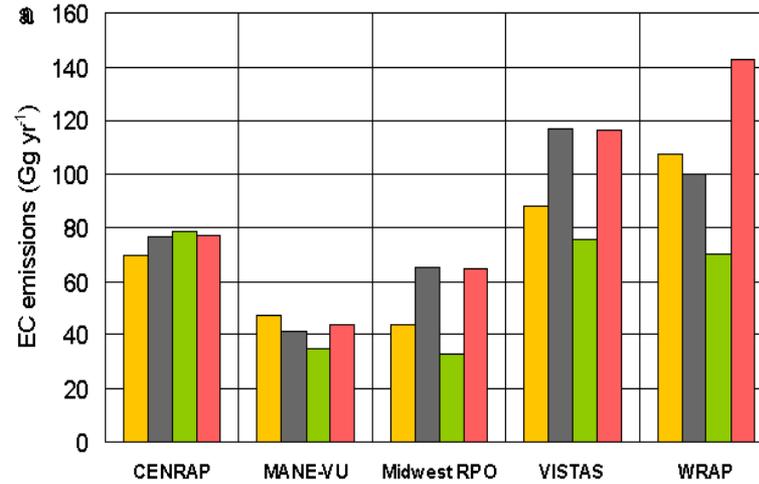
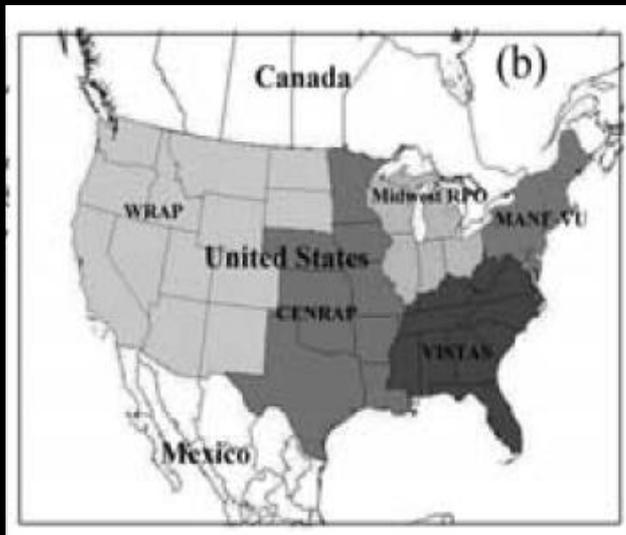
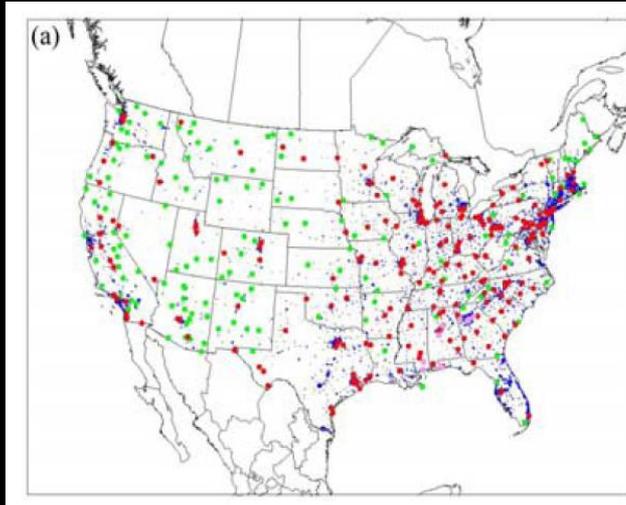
Comparison between Adjoint and Analytical Solution

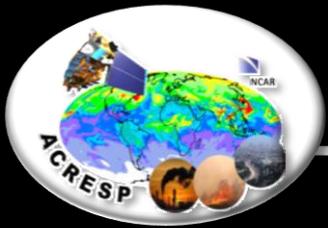




sensitivity to observing network

Top-down Estimates of Black Carbon

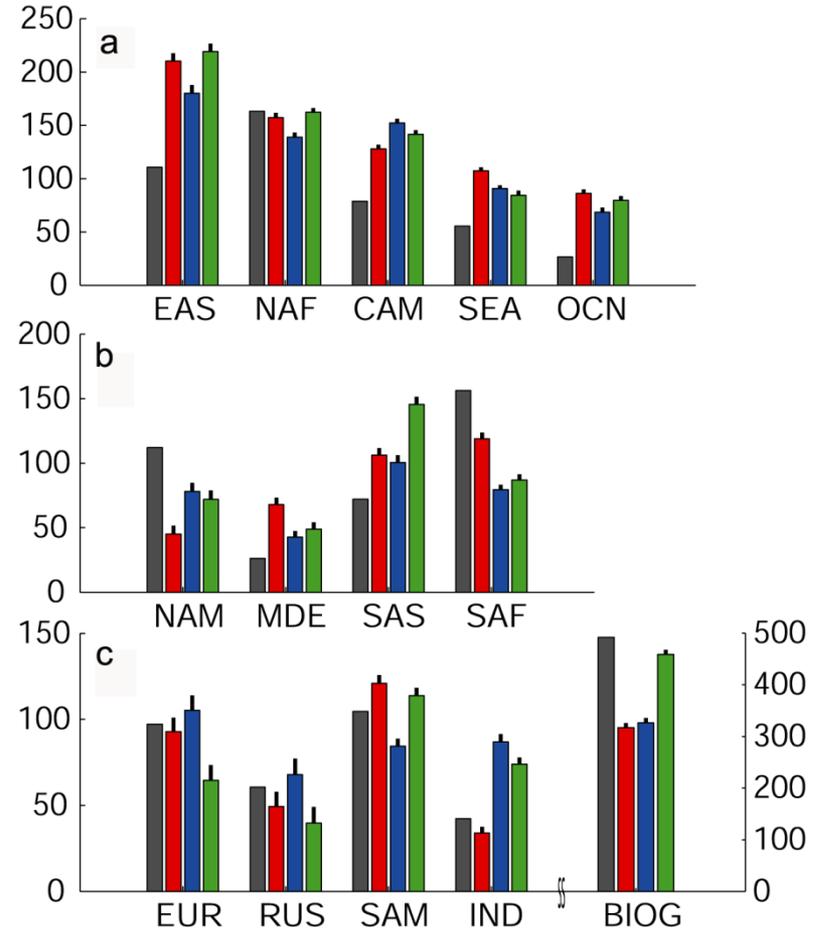
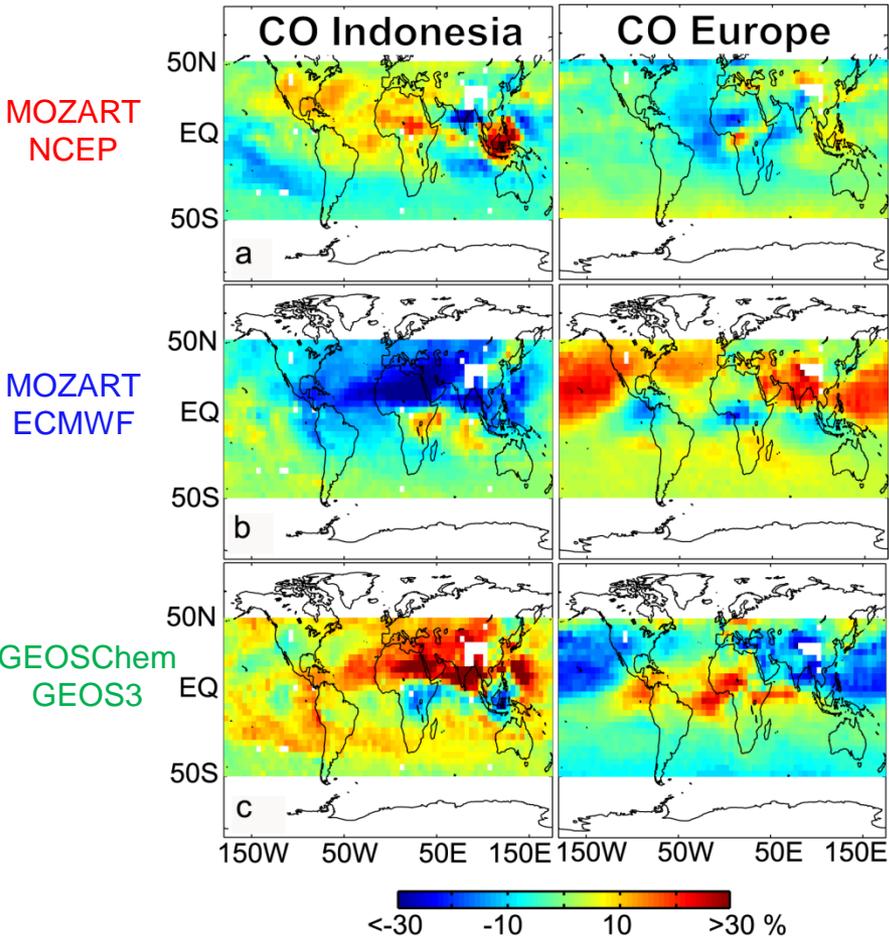


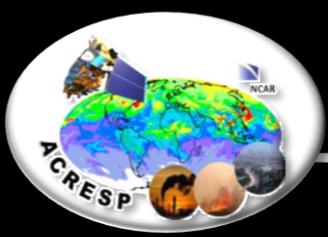


sensitivity to GCTM transport

differences in response functions

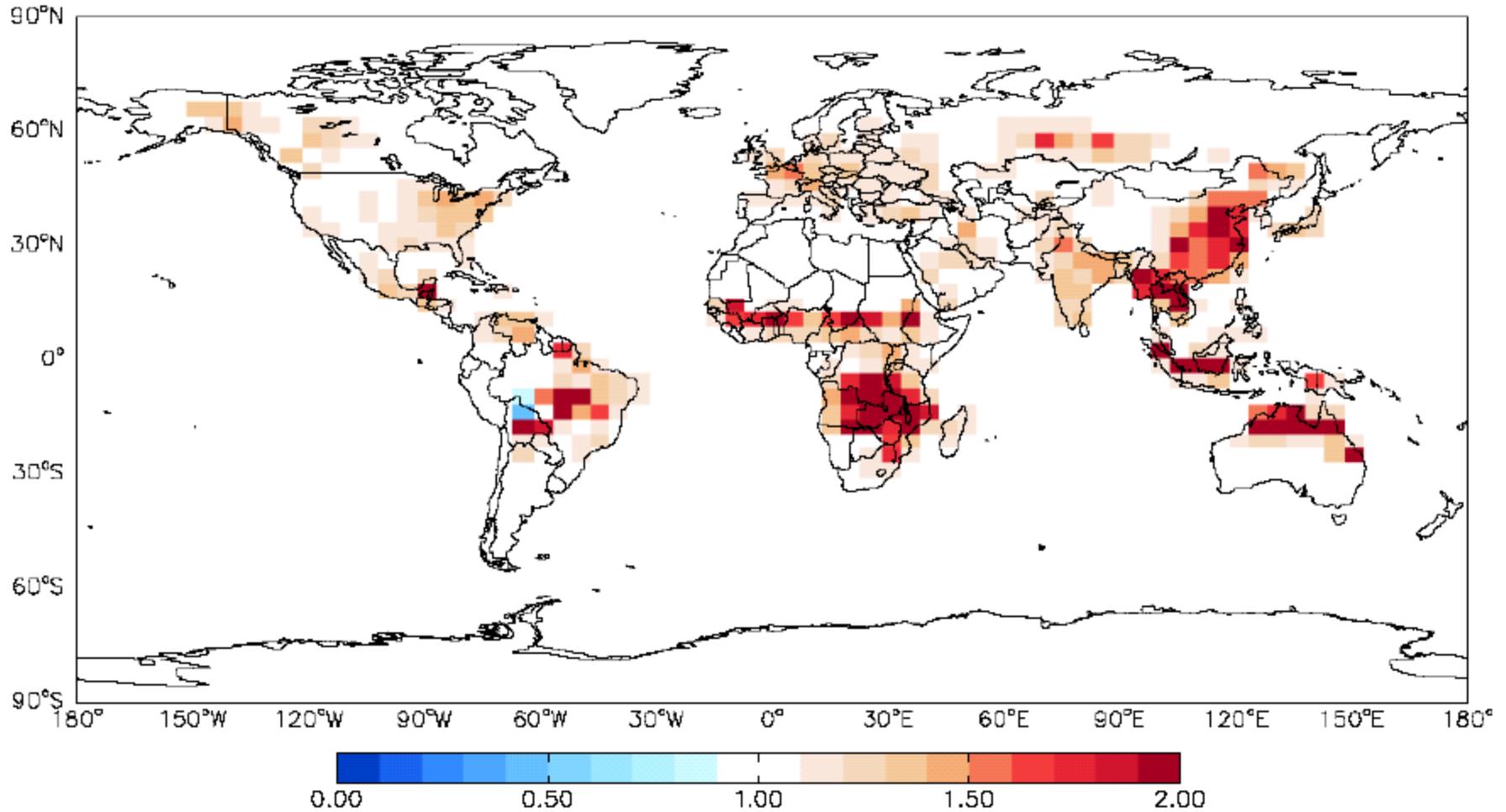
differences in inverse results



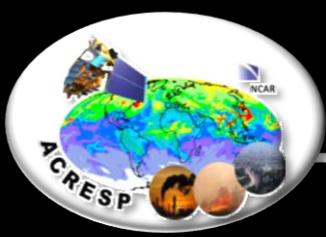


recent top-down estimates of CO sources

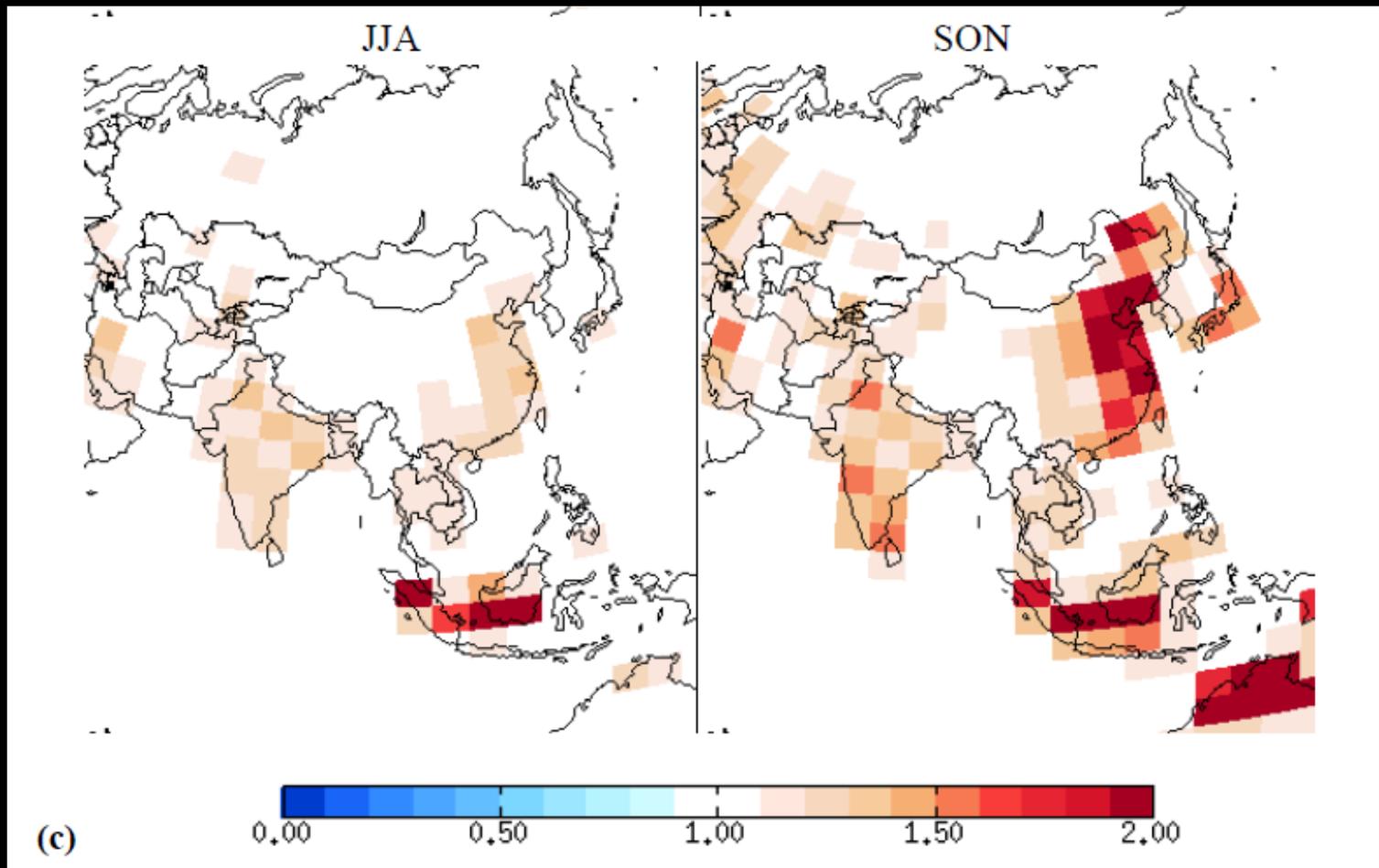
annual estimates of scaling factors

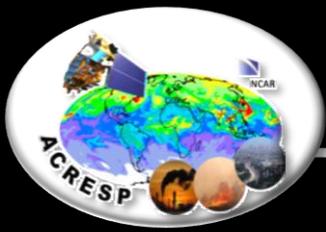


Kopacz et al., (2010)

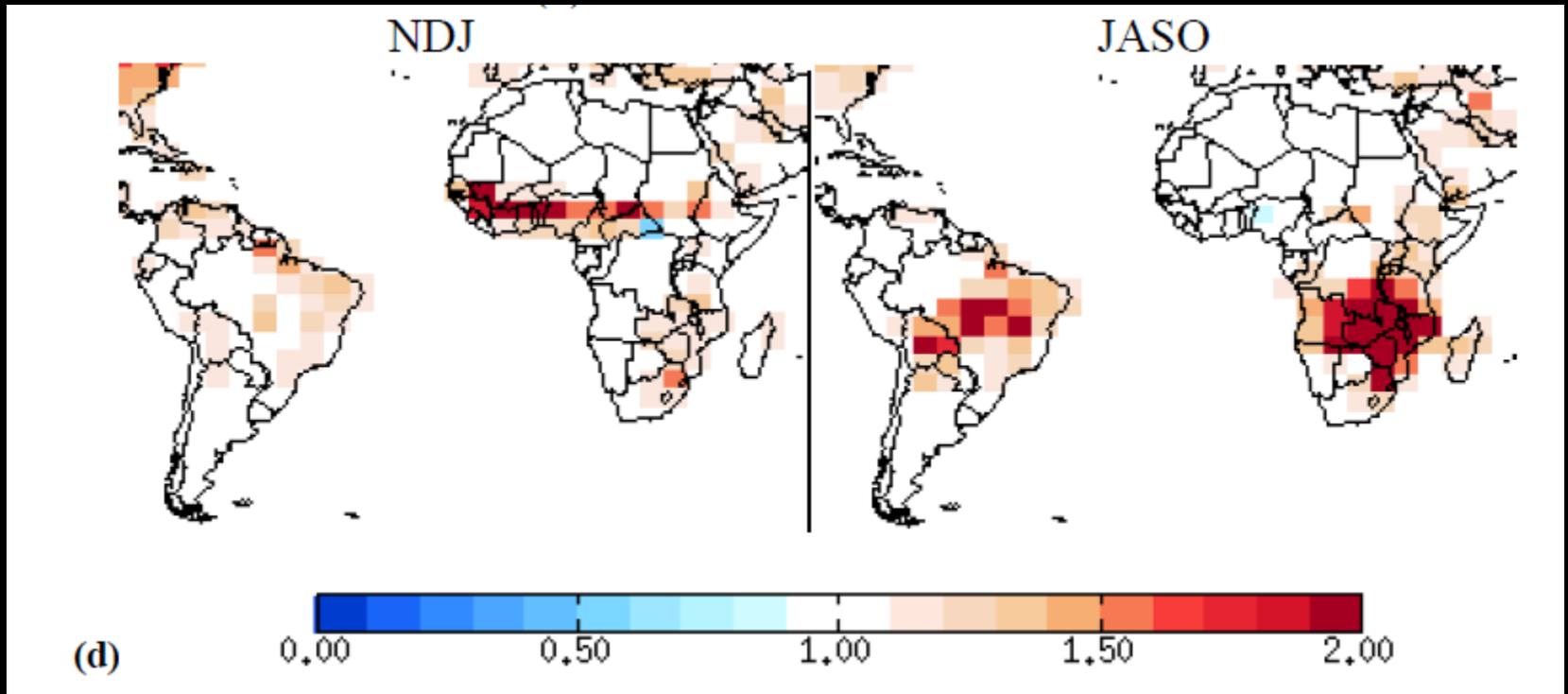


recent top-down estimates of CO sources

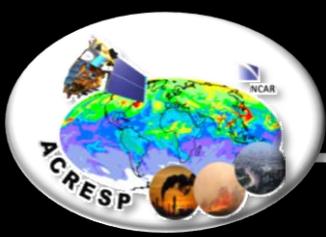




recent top-down estimates of CO sources

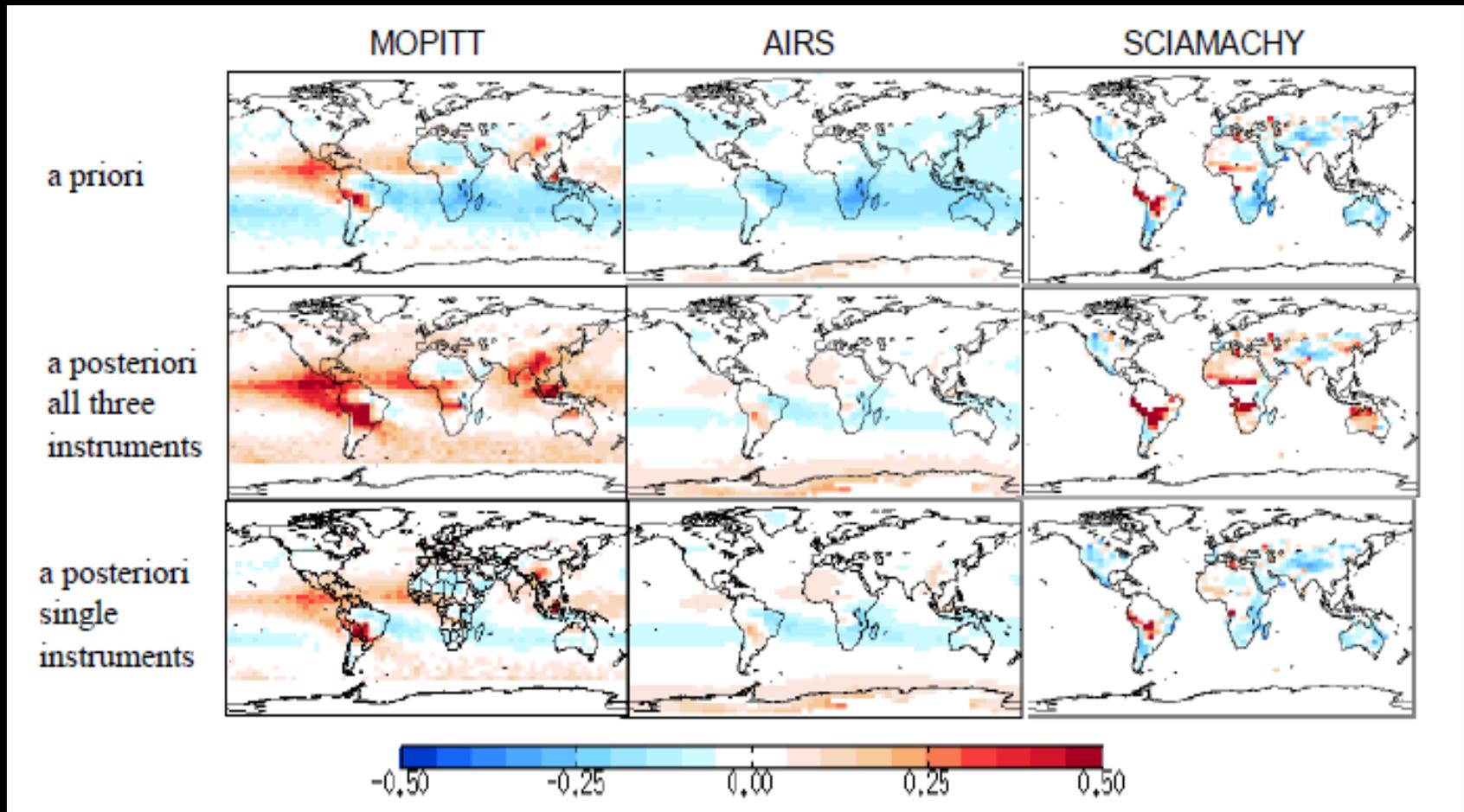


Kopacz et al., (2010)

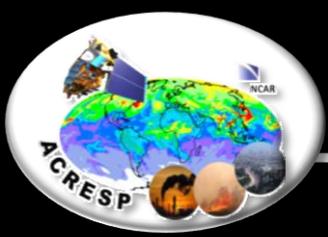


problems?

posterior CO still exhibit large biases

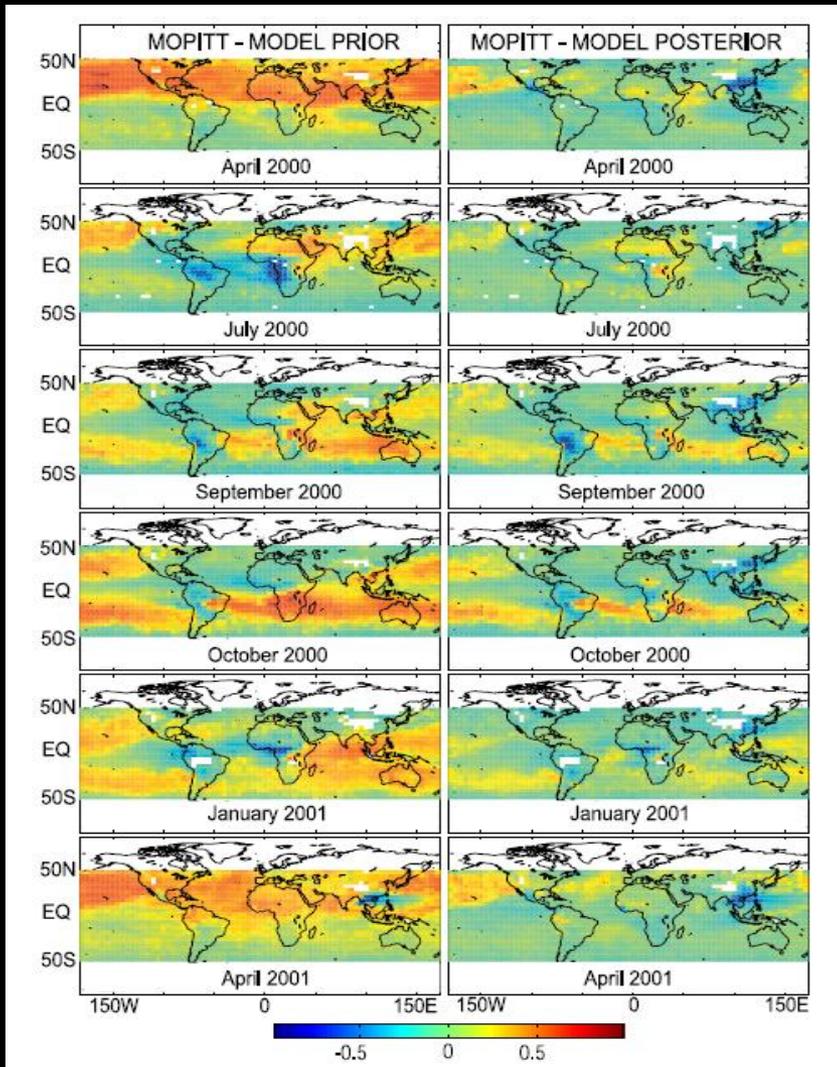


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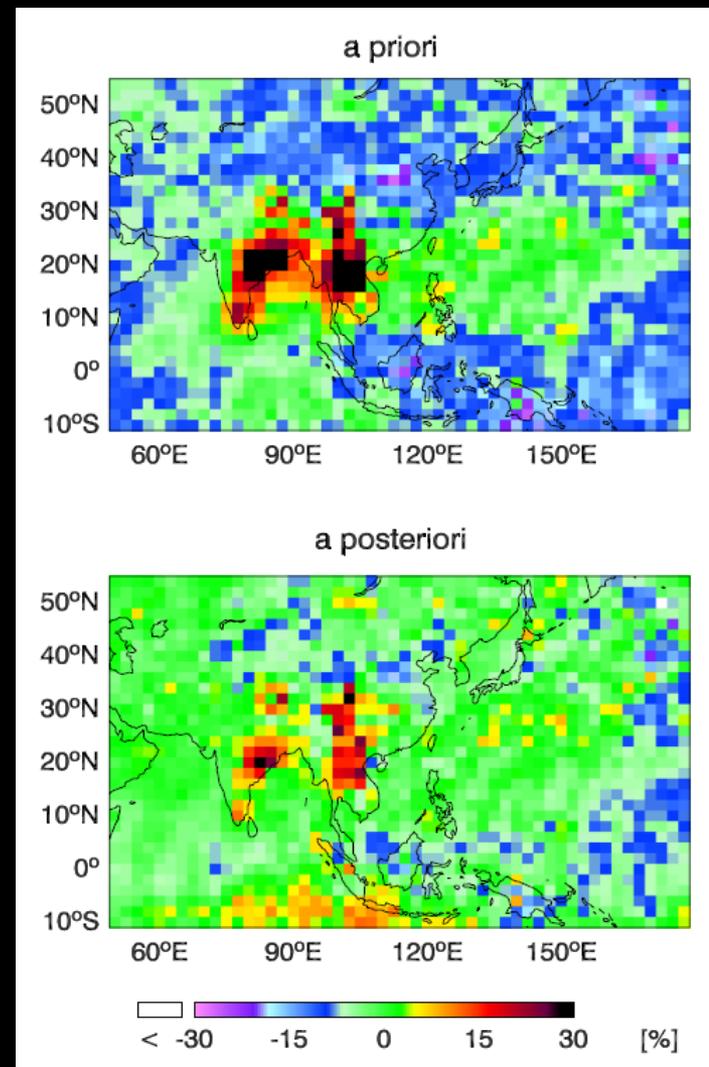


problems?

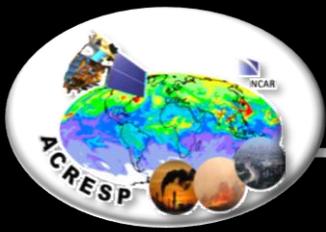
posterior CO still exhibit large biases



Arellano et al. (2006)



Heald et al. (2004)

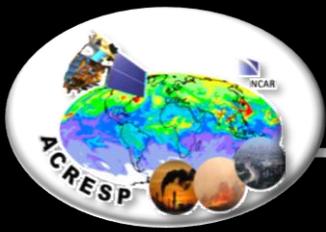


- 1) **Inversions with a regional focus**
 - reconciling differences in inverse results
 - reducing aggregation errors

- 2) **TransCom-like CO experiments**
 - understanding errors in GCTM
 - elucidating differences in methodology (perfect model experiments)

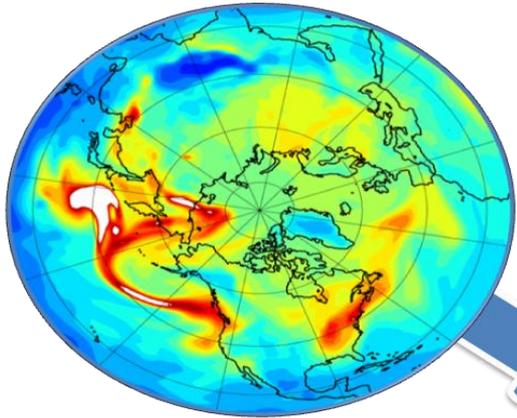
- 3) **Towards a more integrated/synergistic approach**
 - multi-species/multi-platform inversions
 - state/source estimation

- 4) **Better error characterization**
 - observational errors (biases/inconsistencies)
 - model errors
 - emission errors



chemical data assimilation system @ NCAR

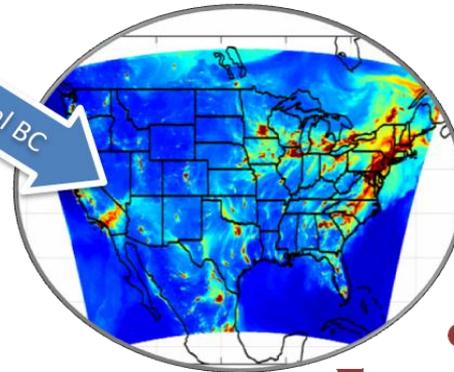
CAM-CHEM/DART



RESEARCH-BASED REGIONAL TO GLOBAL NWP WITH CHEMISTRY

LEVERAGING ON STATE-OF-THE-ART COMMUNITY MODELS (CAM-CHEM, WRF-CHEM) AND COMMUNITY DA FACILITY (DART)

WRF-CHEM/DART



Lateral BC

OBSERVATIONS

CURRENT:

RAWINSONDES

ACARS

AIRCRAFT

SATWINDS

MOPITT CO

MODIS AOD

PLANNED:

TES CO, O₃

IASI CO, O₃

OMI NO₂, O₃

GOME O₃

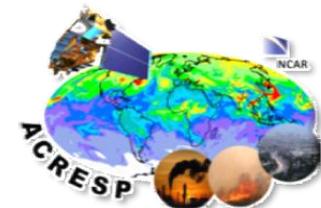
APPLICATIONS

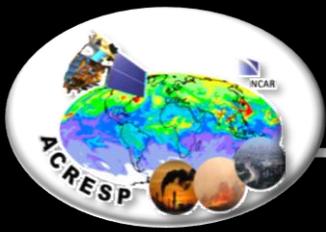
CHEMICAL OSSES

CHEMICAL WEATHER

FIELD CAMPAIGN SUPPORT

IN COLLABORATION WITH NCAR/IMAGE & NCAR/MMM

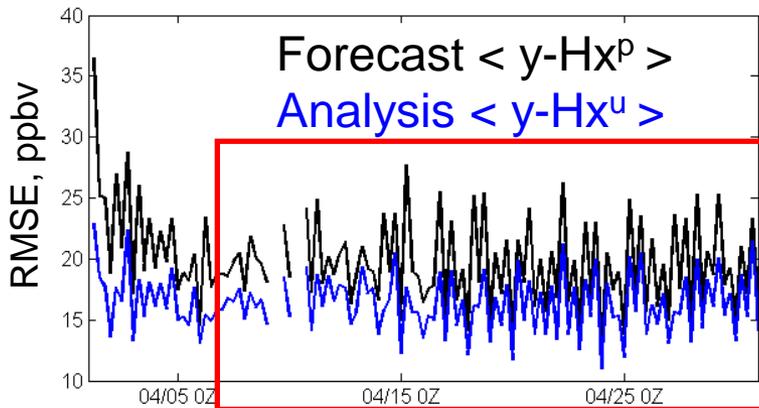




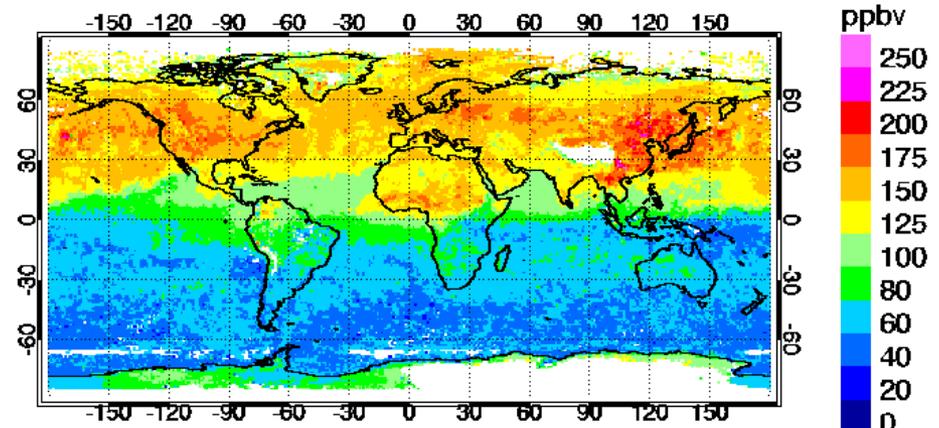
utility of a chemical DA system

capability to evaluate the forecast model

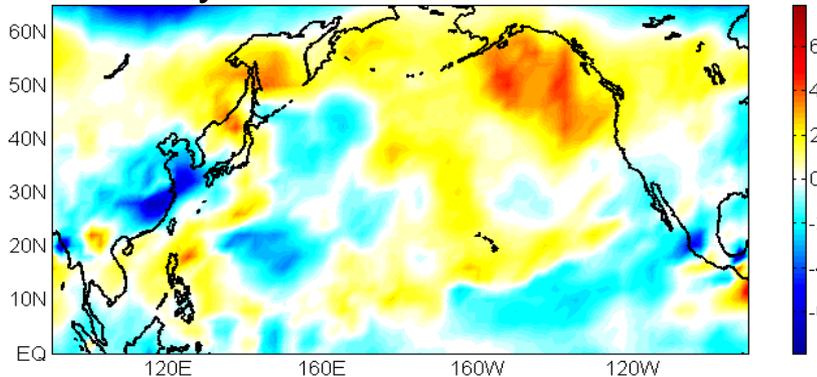
RMSE relative to MOPITT



MOPITT CO @ 700 hPa (April 2006)

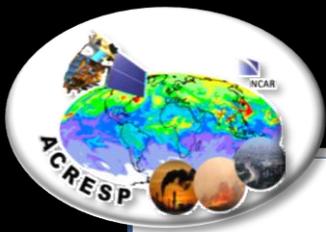


Analysis - Forecast $\langle x^u - x^p \rangle$



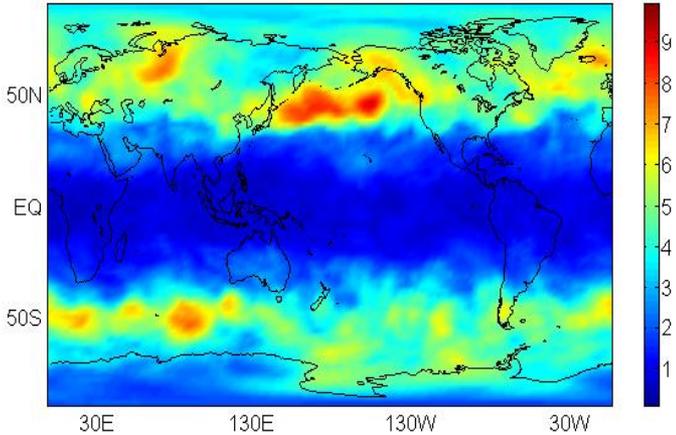
system diagnostic
for short-term
(6-hourly)
forecast errors

model overpredicts
in source region
while underpredicts
in downwind region

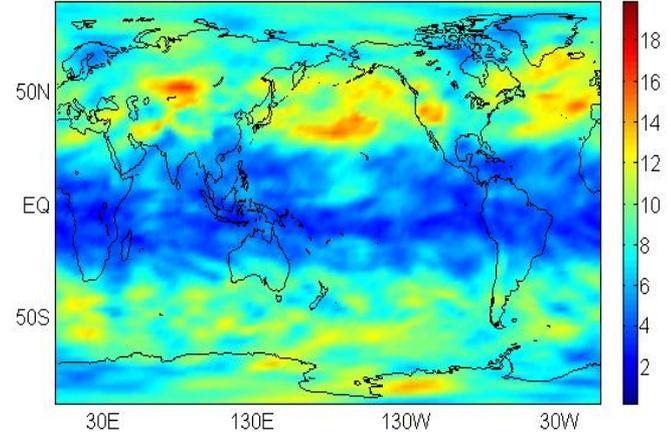


error characterization using ensembles

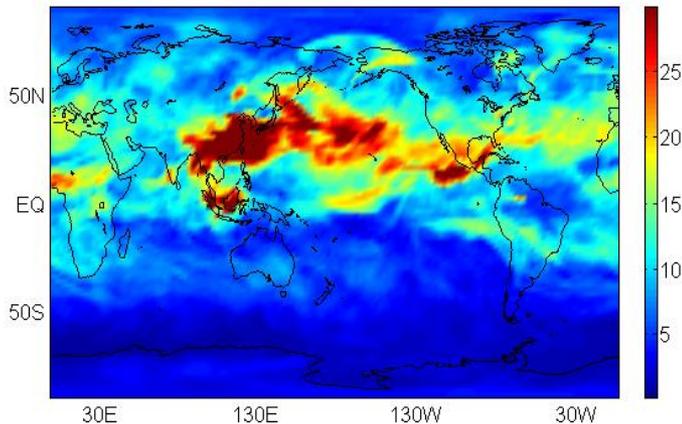
Temperature Spread @ 500 hPa



U Wind Spread @ 500 hPa (m/s)



CO Spread @ 500 hPa (ppbv)



CO Emissions Spread (%)

